2011 GSA Annual Meeting in Minneapolis (9–12 October 2011)

Paper No. 250-6

Presentation Time: 9:00 AM-6:00 PM

AUTOMATIC CONTOURING OF TWO-DIMENSIONAL FINITE STRAIN DATA ON THE UNIT HYPERBOLOID AND THE USE OF HYPERBOLOIDAL STEREOGRAPHIC, EQUAL-AREA AND OTHER PROJECTIONS FOR STRAIN ANALYSIS

VOLLMER, Frederick W., Geology, SUNY New Paltz, New Paltz, NY 12401, vollmerf@newpaltz.edu

Two-dimensional strain data are plotted on a cartesian R_f - ϕ graph (Dunnett, 1969) or a polar strain graph (Elliott, 1970). A logarithmic scale is related to the deviatoric natural strain, ε (Nadai, 1950). Yamaji (2008) showed that a two-dimensional unit hyperboloid H^2 provides a unifying parameter space. Points on H^2 are $\mathbf{x} = (x_0, x_1, x_2)^T$, with origin C. If strain is represented by $(\rho, \psi) = (\ln R, 2\phi)$, then an ellipse is $\mathbf{x} = (\cosh \rho, \sinh \rho \cos \psi, \sinh \rho \sin \psi)^T$. Reynolds (1993) gave equidistant, equal-area, orthographic, gnomic and stereographic azimuthal projections to map H^2 to the x_1x_2 plane. The equidistant is the Elliott graph, it preserves radial distance so ε is undistorted. Wheeler (1984) discussed the orthographic. The equal-area distorts ε but preserves area for comparing densities. The stereographic is conformal. Curves of equal distance from C remain circles with strain, so for a symmetrical distribution the centroid of the projected data is the best-fit ellipse. Projecting H^2 onto a surface whose axis is parallel to x_0 gives a family of cartesian graphs. The equidistant is the R_f - ϕ graph. The centroid in none of these graphs is a good estimator of the best-fit ellipse.

Elliott (1970) hand-contoured strain data on the polar graph to bring out indications of pre-strain fabrics. It is desirable to have a method that is rapid, reproducible, and based on the underlying geometry of the data, rather than the projection. H^2 provides a measure of distance directly related to strain, $d_H = \cosh^{-1}(-\mathbf{a} \cdot \mathbf{b})$, analogous to a great-circle distance on a sphere. By analogy with methods for spherical orientation data (Diggle and Fisher, 1985; Fisher *et al.*, 1987; Vollmer, 1995), contouring strain data can be done by back-projecting a grid onto H^2 using inverse functions. The distances from each node to each data point \mathbf{x}_k are summed to determine the node value, f_{ij} , using a weighting function, w_k , with parameter κ , based on the cumulative distribution function for H^2 (Jensen, 1981). To account for sample size, n, κ is replaced with a normalized parameter: $\kappa_n = \kappa/n^{\frac{1}{3}}$, by analogy with the spherical case (Fisher *et al.*, 1987). The f_{ij} values are contoured as percentages of the maximum f_{ij} value. A computer progam, EllipseFit, that implements these methods is freely available.

2011 GSA Annual Meeting in Minneapolis (9–12 October 2011)
General Information for this Meeting

Session No. 250--Booth# 183

Deformation of the Lithosphere: Field Observations, Experimental Investigations, and Numerical Studies (Posters)

Minneapolis Convention Center: Hall C

9:00 AM-6:00 PM, Wednesday, 12 October 2011

Geological Society of America Abstracts with Programs, Vol. 43, No. 5, p. 605

© Copyright 2011 The Geological Society of America (GSA), all rights reserved. Permission is hereby granted to the author(s) of this abstract to reproduce and distribute it freely, for noncommercial purposes. Permission is hereby granted to any individual scientist to download a single copy of this electronic file and reproduce up to 20 paper copies for noncommercial purposes advancing science and education, including classroom use, providing all reproductions include the complete content shown here, including the author information. All other forms of reproduction and/or transmittal are prohibited without written permission from GSA Copyright Permissions.

1 of 1 10/15/11 6:55 PM